

Combined plate motion and density-driven flow in the asthenosphere beneath Saudi Arabia: Evidence from shear-wave splitting and seismic anisotropy

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ABSTRACT

We analyzed mantle anisotropy along the Red Sea and across the Arabian Peninsula using shear-wave splitting recorded by stations from three different seismic networks—the largest, most widely distributed array of stations examined across the Arabian Peninsula to date. Stations near the Gulf of Aqaba display fast orientations aligned parallel to the Dead Sea transform fault, most likely related to the strike-slip motion between Africa and Arabia. However, most of our observations across Arabia are statistically the same (at a 95 % confidence level), with north-south-oriented fast directions and delay times averaging -1.4 s. Since end-member models of fossilized anisotropy and present-day asthenospheric flow do not adequately explain these observations, we interpret them as a combination of plate- and density-driven flow in the asthenosphere. The combination of northeast-oriented flow associated with absolute plate motion with northwest-oriented flow associated with the channelized Afar upwelling along the Red Sea produces a north-south resultant that matches the observations and supports models of active rifting.

Keywords: Arabia, Red Sea, anisotropy, shear-wave splitting, continental rifting, mantle flow.

INTRODUCTION

When deformed by dislocation creep, olivine in the upper mantle develops lattice preferred orientations (LPO), where the crystallographic a -axes [100] become parallel to the induced shear, leading to velocity variations with propagation direction (Mainprice and Silver, 1993). Shear waves encountering such anisotropic regions split into two orthogonal components, one traveling faster than the other. The anisotropy can be characterized by the polarization direction of the fast wave (cp) and the delay time between the fast and slow waves (δt), and these measurements can be used to provide constraints on the mechanisms causing deformation in the upper mantle (Silver and Chan, 1991; Vinnik et al., 1992). In rift environments, one may expect a rift-perpendicular cp , since LPO should develop parallel to extension for dry upper-mantle conditions (Ribe, 1992). However, in a number of rift environments, such as the Baikal Rift zone, the Rio Grande Rift, and the East African Rift, the observed cp is actually closer to rift-parallel (Gao et al., 1997; Gashawbeza et al., 2004; Walker et al., 2004). This may indicate more complex rifting mechanisms or other deformation processes that dominate ex-

ternally uplifted, with elevation decreasing away from the rift axis, and associated lavas have an alkalic composition, reflecting a deepmantle source. Local convection may lead to more complicated flow patterns and therefore more complex anisotropy at depth. Several studies have suggested that these two end-member models may not be mutually exclusive; rifting in the Red Sea may have been initiated by passive processes, followed by more recent active processes associated with a mantle upwelling (Camp and Roobol, 1992; Ebinger and Sleep, 1998; Daradich et al., 2003).

Several previous anisotropy studies near the Red Sea have revealed fairly consistent patterns. Wolfe et al. (1999) analyzed shear-wave splitting from eight Incorporated Research Institutions for Seismology (IRIS) stations across western Arabia (Fig. 1), which were operated as part of the Program for the Array Seismic Studies of the Continental Lithosphere (PASSCAL), and found St values of 1.0–1.5 s and cp oriented approximately north-south. Receiver function analysis by Levin and Park (2000) found evidence for a more complex anisotropic structure beneath one Global Seismographic Network station, consisting of two dipping layers at depth, but with a resultant cp oriented north-south. Further north, Schmid et al. (2004) and Levin et al. (2006) examined splitting at several stations near the Gulf of Aqaba and the Dead Sea transform fault, where they found average δt of 1.3 s and cp slightly east of north, with some evidence for a more complex, two-layer anisotropic model. However, each of these studies was somewhat limited in their station distribution and data sampling. In this study, we present a more comprehensive analysis of the anisotropic signature along the Red Sea and across Saudi Arabia by analyzing shear-wave splitting recorded by stations from three different seismic networks. This is the largest, most widely distributed ar-

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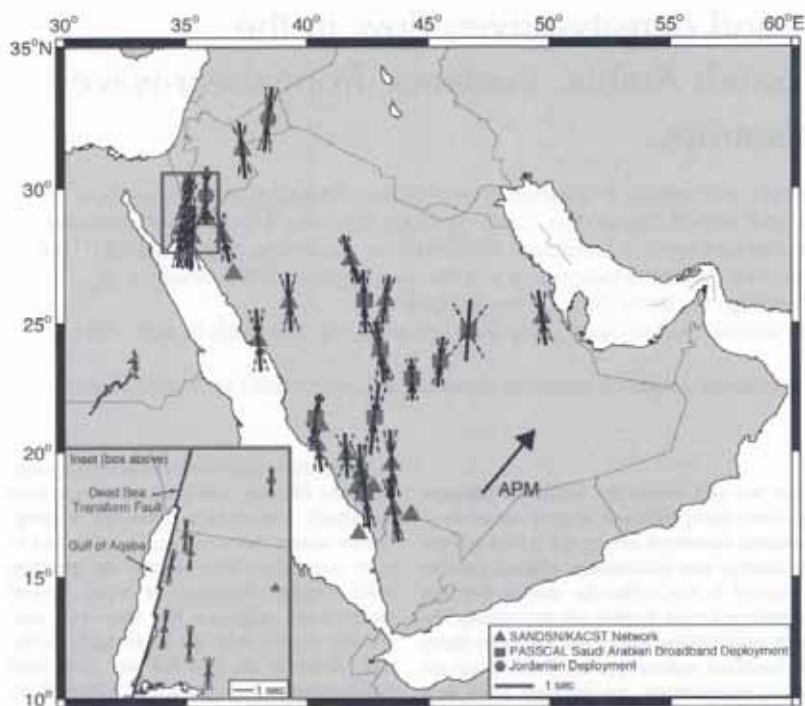


Figure 1. Map showing average splitting parameters. Bold, center lines at each station are oriented in station's average ϕ and length of line is scaled to average δt . Dashed "fans" show one standard deviation of fast angle. Inset provides closer view of Gulf of Aqaba stations (gray box). Triangles—SANDSN stations, squares—PASSCAL stations, circles—Jordanian stations. Black arrow shows average absolute plate motion direction based on global positioning system (GPS) models (Reilinger et al., 1997; McClusky et al., 2003). APM—absolute plate motion.

ray of stations examined across Saudi Arabia to date. We demonstrate that the north-south ϕ is not just valid at isolated sites, but extends throughout the whole of Arabia. These observations cannot be adequately fit by previously proposed models of fossilized anisotropy or present-day absolute plate motion, and we present a new model that consists of a combination of both plate- and density-driven flow in the asthenosphere.

SHEAR-WAVE SPLITTING ANALYSIS AND RESULTS

Teleseismic data recorded on broadband instruments from three different seismic arrays were used. The largest array, the Saudi Arabian National Digital Seismic Network (SANDSN), includes 25 broadband stations distributed along the eastern edge of the Red Sea and across Saudi Arabia (Fig. 1; Al-Amri and Al-Amri, 1999). SANDSN data from events occurring since 2000 were used for this study. To supplement the SANDSN coverage, we also analyzed data recorded by the eight IRIS-PASSCAL Saudi Arabian Broadband Array stations, which operated from November 1995 to March 1997 (Vernon and Berger, 1998), as well as data recorded between 1998

and 2001 from two stations deployed in Jordan (Rodgers et al., 2003; Fig. 1).

We primarily analyzed SKS phases recorded at these stations, but some S and SKKS phases were also included to improve the incidence angle and back-azimuth coverage. The data were bandpass-filtered between 0.05 and 1 Hz to isolate the shear-wave energy, and measurements of the splitting parameters and their associated errors were made using the approach of Silver and Chan (1991). In total, we used 135 events including 247 SKS phases, 12 SKKS phases, and 52 S phases (see the GSA Data Repository material for details¹).

Average splitting parameters from different events at each station are shown in Figure 1. Broadly, all stations display a north-south ϕ with an average δt of 1.4 s, similar to previous findings throughout the area (Wolfe et al., 1999; Levin and Park, 2000). Statistical methods were used to examine the variation in splitting observations at individual stations

¹GSA Data Repository item 2006188, shear-wave splitting analysis, is available online at www.geosociety.org/pubs/ft2006.htm, or on request from editing@geosociety.org or Documents Secretary, GSA, P.O. Box 9140, Boulder, CO 80301, USA.

more closely and to compare the observations from different stations to one another. None of the stations showed significant back-azimuth variation, implying that multiple layers of anisotropy or dipping anisotropic symmetry axes are not required, and the observations at most stations were statistically indistinguishable. Stations near the Gulf of Aqaba (Fig. 1, inset) are a notable exception. While statistically similar to one another, this group of stations displays a statistically different (at a 95% confidence level) average ϕ that is rotated further east than the other stations that we examined.

DISCUSSION

Gulf of Aqaba Stations

Our results in the Gulf of Aqaba region are similar to observations made at nearby stations by Schmid et al. (2004) and Levin et al. (2006). At the two stations they examined, Schmid et al. (2004) found an average ϕ of 3°–8° east of north and a δt of 1.3 s. They argued that ϕ is parallel to the Dead Sea transform fault due to the strike-slip motion between Arabia and Africa. Levin et al. (2006) postulated that their data are better fit by a two-layer model, where neither layer has a ϕ parallel to the Dead Sea transform fault. Their interpretation is that the two-layer model reflects deformation in the asthenosphere caused by absolute plate motion overlain by fossilized anisotropy. Our observations near the Dead Sea transform fault can be fit by a two-layer model similar to the one proposed by Levin et al. (2006); however, this fit is not statistically better than our one-layer model given the greater degrees of freedom. Therefore, we conclude that a one-layer model that has a ϕ parallel to the Dead Sea transform fault can best explain the splitting observations at the Gulf of Aqaba stations. This concurs with the findings of Schmid et al. (2004) and is similar to anisotropy observations made along other transform boundaries (Vinnik et al., 1992; Bostock and Cassidy, 1995).

Stations across Saudi Arabia

Aside from the Gulf of Aqaba stations, most of the splitting observations across Saudi Arabia are very consistent; yet, a straightforward explanation for these observations is difficult to apply. Wolfe et al. (1999) concluded that the predominantly north-south ϕ reflects either fossilized anisotropy or present-day asthenospheric flow. We believe that neither of these end-member models adequately fits the observations. A majority of the examined stations are located on the Arabian Shield, an area composed of Proterozoic terranes that mostly strike north-south (Fig. 2; Stoesser and Camp, 1985). Therefore, the splitting observed at these stations might be attributed to fossilized structure associated with the convergence of these terranes during the Arabian

Figure 2. Topographic map with geologic boundaries. Boundary between Arabian Shield (AS) and Arabian Platform (AP) is marked by solid bold line, and approximate location of Afar upwelling is enclosed by dashed black circle. Thin dashed lines show trend of Najd fault zone, and thin solid lines show boundaries of major terranes, names of which are listed (modified from Stoesser and Camp, 1985).

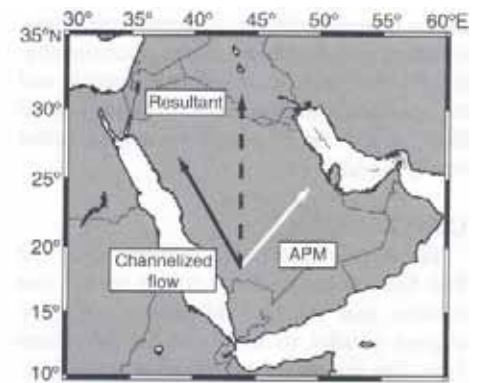


Figure 3. Vector examination of plate motion (white arrow) coupled with channelized upwelling flow (solid black arrow) beneath Saudi Arabia. If we estimate that absolute plate motion is oriented N40°E at a rate of 22 mm/yr (Reilinger et al., 1997; McClusky et al., 2003) and that channelized hotspot flow is oriented approximately N30°W (Camp and Roobol, 1992; Ebinger and Sleep, 1998), then the rate of hotspot flow needed to obtain a north-south resultant (black dashed arrow) is -27 mm/yr. APM-absolute plate motion.

Shield's assembly. However, a 50-100-Myr-period of postaccretionary tectonics produced the northwest-trending strike-slip Najd fault zone, which has had 200-300 km of displacement (Fig. 2; Stoesser and Camp, 1985). Therefore, it seems unlikely that anisotropy imparted during Proterozoic assembly of the Arabian Shield was not disturbed by more recent deformation. For stations on the Arabian Platform (Fig. 2), there is little geologic constraint on any potential fossilized anisotropy because the Proterozoic basement is covered by a thick layer of Phanerozoic sediments (Stoesser and Camp, 1985).

Fossilized anisotropy can also be constrained by lithospheric thickness. Using common percentages of anisotropy (4%-5%; Mainprice and Silver, 1993), the lithosphere beneath Saudi Arabia would need to be ~140 km thick to accumulate the observed δt . Although sparse in this region, estimates of lithospheric thickness have been determined with seismic refraction and receiver function methods and are 100-120 km near the Arabian Shield-Platform boundary (Mooney et al., 1985) and 80-100 km within the Arabian Shield (Sandvol et al., 1998). Closer to the Red Sea, xenolith and isotope studies have indicated that the lithosphere thins to as little as 40 km (Altherr et al., 1990; Camp and Roobol, 1992). This substantial thinning is corroborated by seismic waveform modeling, P-wave tomography, and Sn attenuation studies in western Arabia (Rodgers et al., 1999; Benoit et al., 2003; Al-Damegh et al., 2004). Therefore, the lithosphere is probably not thick enough, especially near the Red Sea Rift, to generate the δt observed. In addition, because the lithosphere thins from the Arabian Shield to the Red Sea, fossilized anisotropy should produce a pattern of increasing δt at stations from west to east. No such pattern exists, and instead our δt values across Saudi

Arabia are remarkably consistent (Fig. 1), casting further doubt on fossilized anisotropy as a viable model for this region.

An alternative explanation for the observed splitting is the alignment of cp with absolute plate motion due to shear at the base of the lithosphere. Several estimates of the absolute plate motion for Arabia have been proposed using a variety of global plate models. Gripp and Gordon's (2002) HS3-NUVELIA model produces an absolute plate motion direction for Arabia of about N55°W. However, their model is based only on Western Hemisphere hotspots and does not fit African hotspots well. Therefore, this model probably does not provide the best estimate for Arabian absolute plate motion. Instead, global positioning system (GPS) plate models indicate that Arabia is moving northeast, with an average rate and direction of 22 mm/yr and N40°E, respectively (Reilinger et al., 1997; Sella et al., 2002; McClusky et al., 2003). The GPS results are comparable to Wolfe et al.'s (1999) absolute plate motion calculation, based on the O'Conner and le Roex (1992) African absolute plate motion model. When compared to the GPS-determined absolute plate motion direction, the orientation of cp is at least 30° different; therefore, this explanation also fails to fit the observations.

Since these end-member models do not fit the splitting observations, we conclude that the anisotropy is the result of an interaction of mantle flow in the asthenosphere. Shear caused by absolute plate motion, directed approximately N40°E at 22 mm/yr (Reilinger et al., 1997; McClusky et al., 2003), may affect the alignment of mantle minerals. However, based on variations in the topography and the distribution of alkalic volcanics, it has also been suggested that flow radiating from the Afar mantle upwelling is channelized toward the Red Sea Rift (Camp and Roobol, 1992;

Ebinger and Sleep, 1998), in a direction oriented at about N30°W. Assuming that the strain caused by this upwelling is comparable to that caused by plate motion, we can combine these two flow orientations, similar to the vector approach of Silver and Holt (2002), to obtain a north-south-oriented resultant (Fig. 3).

Several other lines of evidence also help support our conclusions. Seismic tomography models show that the upper mantle beneath western Arabia is anomalously slow, with velocities increasing toward the continental interior (Debayle et al., 2001; Benoit et al., 2003). These observations are attributed to thermal differences and indicate much hotter mantle beneath the Red Sea, which is consistent with flow directed beneath the rift. Daradich et al. (2003) demonstrated that the higher elevations along the Red Sea and the overall tilt of the Arabian plate result from viscous stresses associated with large-scale mantle flow from the Afar upwelling. Walker et al. (2005) compared splitting results from both Wolfe et al. (1999) in Saudi Arabia and Gashawbeza et al. (2004) in Ethiopia to the parabolic asthenospheric flow model expected from a plate moving over a stationary hotspot. Over 90% of the splitting observations are consistent with radial flow, and they attributed the remaining variations to flow-modifying factors, such as channelization by topography at the base of the lithosphere. In addition, Schilling et al. (1992) found isotopic evidence for mantle mixing between depleted asthenosphere and plume flow in Saudi Arabia, supporting the idea of flow interaction at depth. This combination of both plate- and density

driven flow explains the consistent anisotropic signature across Saudi Arabia. Additionally, the rift-oblique cp, the higher topography, and the presence of alkalic lava extrusions near the Red Sea Rift are all consistent with an active rifting model.

CONCLUSIONS

Teleseismic shear-wave splitting along the Red Sea and across Saudi Arabia reveals that stations near the Gulf of Aqaba display cp aligned parallel to the Dead Sea transform fault. The remaining observations across Saudi Arabia show a consistent north-south cp with 8t averaging 1.4 s. Present-day plate motion does not match the observed cp orientation, and lithospheric thickness constraints indicate that fossilized anisotropy cannot explain the observed 8t. Therefore, we interpret the anisotropic signature as a combination of plate- and density-driven flow in the asthenosphere. A combination of the northeast-oriented flow associated with absolute plate motion with the northwest-oriented flow associated with the channelized Afar upwelling generates a north-south-oriented resultant that matches our splitting observations. Other evidence supporting channelized flow is also consistent with active rifting processes.

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