

Local and Regional Earthquake MD Calibration of Tabuk Sub-network, NW of Saudi Arabia

By

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ABSTRACT

The Seismic Studies Center at the King Saud University has established a seismic sub-network at the NW of Saudi Arabia in 1985. It was named the Tabuk sub-network, after the city of Tabuk in the area. A data set of 194 earthquakes were selected and their corresponding ML values calculated by the nearest seismic network (ISSN). 137 observations of these earthquakes were gathered from AYN station, 82 observations from BADA station, 162 observations from HQL station and 65 observations from SRFA station. Four steps analytical procedure were carried out as follows: a) Compilation of data for multi-linear regression, b) Linear regression of the common logarithm of the base 10 of the signal duration (τ), c) Smoothing of the data, and d) Multiple linear regressions of the main variables. The following range of values was incorporated for empirical constant's a_1 , a_2 , a_3 . a_1 vary from -3.05 to -1.68, a_2 vary from 2.17 to 2.61 and a_3 (Which is the epicentral distance correction) vary from 0.003 to 0.004, while the standard deviations vary from 0.17 to 0.22 and the correlation coefficients vary from 0.943 to 0.970. The empirical constants are determined by stepwise multiple linear regression. The local magnitude was assigned as the dependent variable, while the common logarithm of the signal duration, epicentral distance, and the focal depth were assigned as the independent variables. The reasons of using MD in this sub-network are as follows: a) it does not require any amplitude calibration; b) It is not affected by instrument saturation; c) the often-negligible distance-dependence makes the method applicable and d) this method has an extensive application, especially in micro-earthquake surveys.

KEY WORD

Seismic station, Duration magnitude, Local magnitude, Duration, Saudi Arabia, Seismic network, Seismic sub-network.

INTRODUCTION

Magnitude is a term normally used to describe the size of an earthquake based on instrument measurements. Richter (1935) first introduced it. He proposed a magnitude scale based solely on amplitudes of ground motion recorded by seismographs. Richter's procedure to estimate earthquake magnitude followed the practice by Wadati (1931) in which the calculated ground amplitudes for various Japanese stations were plotted against their epicentral distances. Richter (1935) took several bold steps to simplify the estimation of earthquake magnitude.

The magnitude is a quantity, which is a characteristic of an earthquake and can be determined from different recording stations within accepted error limits. Even in the best conditions, errors of 0.2 - 0.3 units are expected in a calculated magnitude. So far, the magnitude concept appears simple and relatively free from complications. However, this is not the case, as evidenced by the three different scales commonly in use: Local magnitude (ML) is the original magnitude introduced by Richter (1935) for local earthquakes in Southern California. *ML* was defined by Richter as the logarithm of the maximum recorded trace amplitude (expressed in microns) by a Wood-Anderson Torsion Seismograph at an epicentral distance of 100 km with specified constants. Gutenberg (1945) developed the concept of a magnitude scale (later on called Surface Wave Magnitude (Ms)) that was applicable to any epicentral distance and to any type of seismograph. This magnitude required a better knowledge of the variation of wave amplitudes with distance. It was necessary to use ground amplitudes instead of trace amplitudes in order to use different seismograph types. The first generalization was made for Rayleigh surface waves, by limiting the period range considered to 20 ± 2.0 sec. Further, by including only earthquakes of normal depth (constant h) and finally by stating that the function f for surface waves is proportional to the logarithm of the epicentral distance. The amplitude refers to the horizontal component of Rayleigh surface waves, and similar formulas for the vertical component of Rayleigh surface waves have been developed. Gutenberg (1945) extended magnitude determinations to body waves (later on called Body Wave Magnitude (mb)) and to earthquakes at any depth.

However, magnitudes given by the three scales do not agree with each other, which mean the same earthquake will have different magnitude values on the different scales.

Magnitude can be calculated from the maximum amplitude on the seismogram (Eaton, et al. 1970) or from the length of time from the P-wave arrival to the end of coda (signal duration). Lee et al. (1972) established an empirical formula for estimating magnitudes of local earthquakes using signal durations.

However, the main purpose of most micro-earthquake networks, which operate short period, high-gain instruments, is to locate every small event, using the arrival times of P-waves and S-waves. In these networks, it is often impracticable to establish a suitable magnitude scale based on amplitude measurements because of the instrument's limited dynamic range, which often results in even relatively small events overloading the recorders.

Tabuk seismographic Sub-Network operates analogue basic equipment (Sinno et al., 1988; Al-Arifi and Gharib, 1994) so this necessitates the establishment of a simple and easy-to-apply magnitude scale, which has none of the problems of the recorder. It consists of 8 seismic stations (figure 1), namely, Haql (HQL), Al-Ouyannah (AYN), Al-Sharaf (SRFA), Al-Sultania (SALT), Bir-El-Mashi (BMSH), Al-Bada (BADA), and Maqna (MKNA). See table 1 for the information pertaining to the geographical location, means of data relay and date of operation of the Tabuk sub-network, Saudi Arabia.

DURATION MAGNITUDE

Estimating magnitudes from signal duration is an empirical correlation between the magnitudes of the earthquakes and the duration of its recorded signal. This correlation has been consistently observed in different parts of the world. A partially satisfactory theoretical basis for this correlation has been sought in terms of the properties of coda waves (Suteau and Whitcomb 1979; Lee and Stewart 1981).

Bisztricsany (1958) was the first to use signal duration in estimating magnitudes by establishing a relationship between the magnitudes of teleseismic events at epicentral distances between 4° to 160° recorded in eastern Europe (Budapest, Prague and Warsaw) and the logarithm of the duration of their surface wave train. He suggested that:

$$M = 2.92 + 2.25 \log(\tau) - 0.001\Delta^\circ \quad (1)$$

Where:

M is magnitude (mainly in the range of 5 to 8).

τ is the duration of the surface wave in seconds.

Δ is the epicentral distance in degrees.

Solov'ev (1965) and Tsumura (1967) derived empirical relationships between the total signal duration of local earthquakes and magnitude for the areas of Sakhalin, USSR, and Kii Peninsula, Japan, respectively. The Tsumura empirical formula was:

$$M = -2.53 + 2.85 \log (F-P) + 0.0014\Delta \quad (2)$$

Where:

M is the magnitude (mainly in the range of 3 to 5).

F-P is the total duration of oscillation in seconds.

Δ is the epicentral distance in kilometers.

The studies of Lee et al. (1972) for central California, USA, and Crosson (1972) for the Puget Sound region of Washington State, USA, have subsequently led to the widespread use of signal duration in estimating magnitudes (for a list of examples, see Adams, 1977; Lee and Wetmiller, 1978; Båth, 1981).

METHOD AND RESULT

According to Lee and Stewart (1981), the duration magnitude (MD) is usually given in the form:

$$MD = a_0 + a_1 \log \tau + a_2 \Delta + a_3 h \quad (3)$$

Where

τ = signal duration in seconds,

Δ = epicentral distance in km,

h = focal depth in km,

a_0, a_1, a_2, a_3 = empirical constants.

The coefficients in equation 3 are usually determined by correlating signal duration for a chosen number of events with their magnitudes, using some other accepted magnitude

scale taking into account the epicentral distance and focal depth. Since no dependence of duration magnitude on depth (range 0.0 - 60.0 km) has ever been conclusively found anywhere, the depth dependent term has been omitted. Equation 3 is therefore reduced to:

$$MD = a_0 + a_1 \log \tau + a_2 \Delta \quad (4)$$

Although IASPEI's commission on practice defined the signal duration as "*The time in seconds between the first onset and the time that trace never again exceeds twice the noise level which existed immediately prior to the first onset*" (Lee and Stewart, 1981), many authors have defined it in different ways to suit their situation.

In this study, signal duration is arbitrarily defined as the time interval in seconds between the first onset of the first P-wave and the point where the signal (peak-to-peak amplitude) no longer exceeds 2 mm as it appears on a TELEDYNE GEOTECH helicorder model RV301B, Portacorder model RV-320B. This is for a recording system consisting of TELEDYNE GEOTECH S-13 Seismometer (Natural period = 1.0 sec/cycle), amp-mod (gain \approx 84 dB), Portacorder (gain \approx 54 dB), and Helicorder (gain =24 dB). In figure 2 signal duration (τ) as defined in this study, with background noise (approximately equal to 1.0 mm) on the portacorder records.

There are three reasons why total duration is defined in this study:

a) This method enables signal duration to be defined independent of the particular recording device,

b) neither of the seismic stations in this study operate at unusually high gain which would tend to make the total duration inconveniently long, or the amplitude of the background noise excessively large, presenting practical difficulties in the measurement of duration, and

c) The network is large in aperture, and, hence, for small earthquakes, the more distant stations may record very small amplitudes (<2.0 mm), which makes it difficult to assign an arbitrary cut-off amplitude but, on the other hand, more practical to read the entire duration. Another point of interest is that defining the end of the record in terms of the background noise makes the duration somewhat independent of a subsequent change in the instrument gain since increasing the magnification would amplify the noise as well as the

earthquake signal.

There are also four reasons why duration magnitude is used in this study:

- a) It does not require any amplitude calibration of the seismograph,
- b) It is not affected by instrument saturation (record clipping) or too faint traces at the largest amplitudes,
- c) The often negligible distance-dependence makes the method applicable also when the epicentral locations are only approximate, and
- d) This method has an extensive application, especially in micro-earthquake surveys, notably in the U.S.A since around 1972.

This method may break down when a new shock occurs on the tail of a preceding event.

One hundred and ninety four earthquakes (506 observations) recorded from January 1986 to December 1991 by the Tabuk Sub-network (Table 1) were sampled to study the duration magnitude. Signal durations of 104, 98, 60 and 44 selected earthquakes recorded by HQL, AYN, BADA and SRFA seismic stations respectively are plotted against their ML (Figures 3, 4, 5, 6).

If the duration magnitude is to be both practical and reliable, it is essential that data be sampled over a wide range of magnitude. The highest magnitude used in this study (4.8) was limited by practical difficulties encountered in reading the duration of large earthquakes, which were almost invariably followed by immediate aftershocks.

The reduction of data was accomplished by a sequence of steps that allowed an examination of the effects of the secondary variables at various stages of the process. The analytical procedure may be outlined as follows:

a) Compile data for multi-linear regression:

A data set of 194 earthquakes (recorded by the most continuous recording stations, AYN, BADA, HQL, and SRFA) were selected and their corresponding ML values calculated by the nearest seismic network (ISSN). 137 observations for these earthquakes were gathered from AYN station, 82 observations were gathered from BADA station, 162 observations were gathered from HQL station and 65 observations

were gathered from SRFA station. The earthquake data were compiled into a data matrix for each station. Epicentral distances and focal depths were calculated by HYPOELLIPSE (Lahr, 1990). The error is approximately 0.2% per kilometer for events up to 400 km in epicentral distance. In view of the small coefficients obtained for the distance term in previous studies, it is indicated that the error involved is negligible in this study. The depth data were included in the above matrix to confirm that there are no effects of the focal depth in this study.

b) Linear regression of the common logarithm of the base 10 of the signal duration (τ):

In this step, the common logarithm of the signal durations gathered from each station were tested, using a simple linear regression as an independent factor and the Local Magnitude as, dependent factor, using a statistical software package called 'SPSS'.

In this step, also abnormal values of signal duration were discarded (longer than expected or shorter than expected). The correlation coefficients (R) are 0.939, 0.864, 0.862, and 0.905 for SRFA, HQL, BADA, and AYN stations respectively.

c) Smoothing of Data:

In this step, the predicted value is calculated and those data that give the higher residual are discarded.

d) Multiple linear regressions of main variables:

Local magnitude determined from short-period seismographs is linearly dependent on the common logarithm of the signal duration and epicentral distance (Lee et al., 1972; Tsumura, 1967). Figures 3, 4, 5, and 6 show the common logarithm of duration for each of the stations studied, plotted against the local Richter magnitude. The linear relationship is evident. All the four stations show a slight increase of slope with increasing magnitude. Based on these figures and previous studies, the previously mentioned linear model (equation 3) is initially assumed to be the most appropriate.

geology.

The adjusted R square values show that 91.26%, 88.53%, 90.25%, and 93.81% of the variations of the local magnitude are explained by the effect of the three independent variables ($\log \tau$, Δ , and focal depth) of the AYN, BADA, HQL, and SRFA stations respectively. The remaining percentages (8.74%, 11.47%, 9.75%, and 6.19%) are due to error, while the F-test shows that the reliability of the cause and effect between the dependent variable and the independent variables are highly significant $F= 507.50, 228.90, 477.94,$ and 327.02 for the AYN, BADA, HQL and SRFA stations respectively. Furthermore, these significant F's values show that the probability of rejecting such relationships is negligible. However, by applying the t-test which defines the degree of reliability between each of the independent variables and the dependent variable, the following results emerged:

- a) the common logarithm of the signal duration ($\log \tau$) t-test shows that it is the most reliable result since the t-test values are 28.18, 18.78, 25.52, and 22.72 for the AYN, BADA, HQL and SRFA stations respectively, with a negligible probability of rejecting this relationship (0.0 for all stations).
- b) Epicentral distance (Δ) t-test shows that it gives less reliable results since the t-test values are 9.99, 8.85, 12.71, and 6.30 for the AYN, BADA, HQL and SRFA stations respectively, with a negligible probability of rejecting this relationship (0.0 for all stations).
- c) Focal depth (h) t-test shows that it produces results of very low reliability since the t-test values are -0.936, -0.496, 0.873, and -1.490 for the AYN, BADA, HQL, and SRFA stations respectively, with a high probability of rejecting this relationship (35.18%, 62.19%, 38.49%, 14.41%, for the AYN, BADA, HQL, and SRFA stations respectively). Therefore, this variable was not included in the equation. Then the magnitude equations were written for the four seismic stations (equation 5 for AYN, equation 6 for BADA, equation 7 for HQL, equation 8 for SRFA) as follows:

$$MD_{\text{AYN}} = -3.01 + 2.57\log \tau + 0.003 \Delta \quad (5)$$

$$MD_{\text{BADA}} = -3.05 + 2.61\log \tau + 0.004 \Delta \quad (6)$$

$$MD_{\text{HQL}} = -1.92 + 2.17\log \tau + 0.004 \Delta \quad (7)$$

$$MD_{SRFA} = -1.68 + 2.19 \log \tau + 0.003 \Delta \quad (8)$$

Table 2. Summary of regression results.

Station	a ₁	a ₂	a ₃	S.E.	R	N
AYN	-3.01±0.37	2.57±0.18	0.003±0.0007	0.21	0.95625	98
BADA	-3.05±0.55	2.61±0.28	0.004±0.0008	0.22	0.94302	60
HQL	-1.92±0.33	2.17±0.17	0.004±0.0006	0.20	0.95102	104
SRFA	-1.68±0.37	2.19±0.19	0.003±0.0008	0.17	0.97006	44

S.E. Standard Error

R multiple correlation coefficient

N Number of events used in the regression

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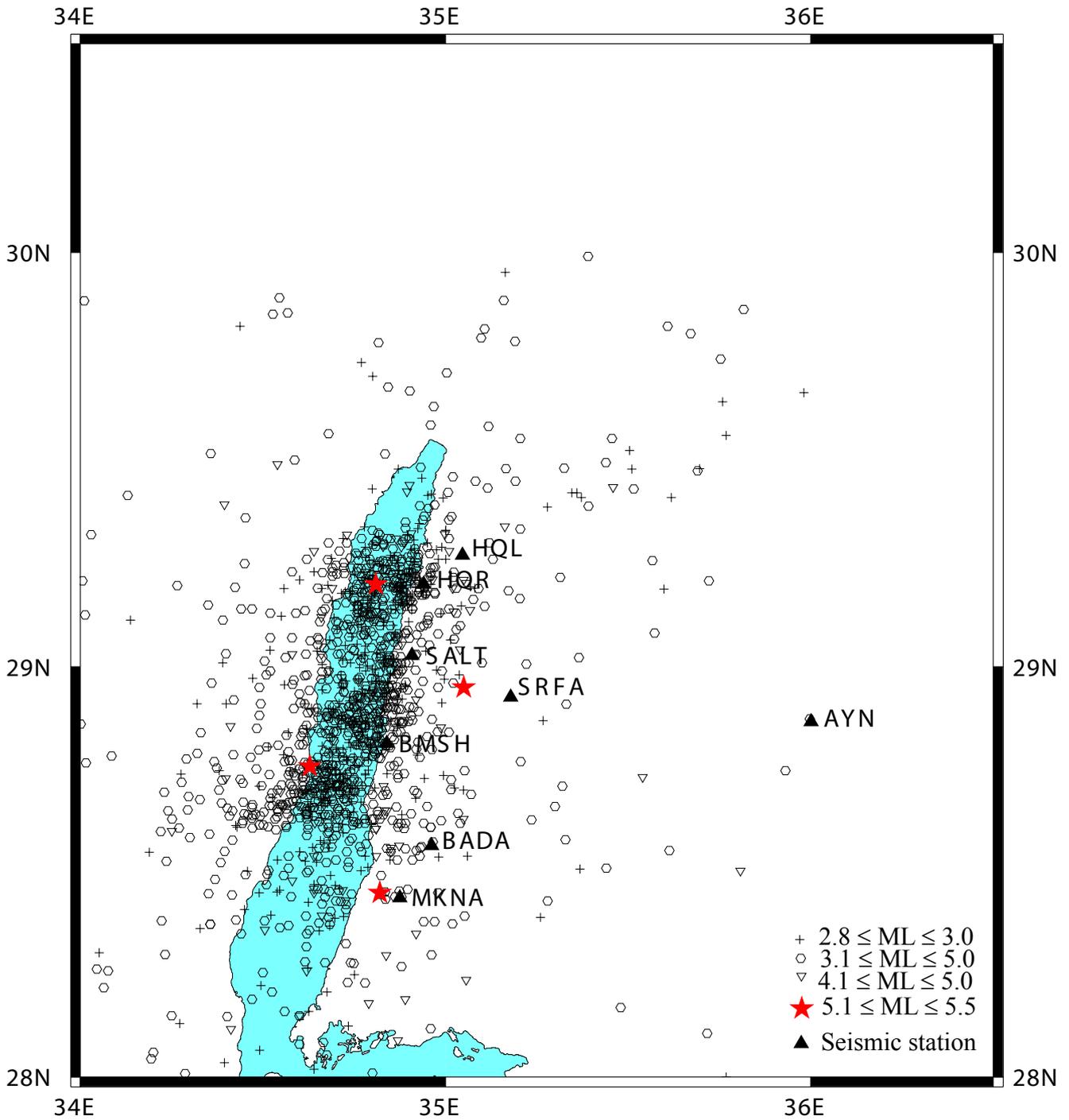


Figure 1 map showing the seismicity of the study area and the location of each seismic station used in this study.

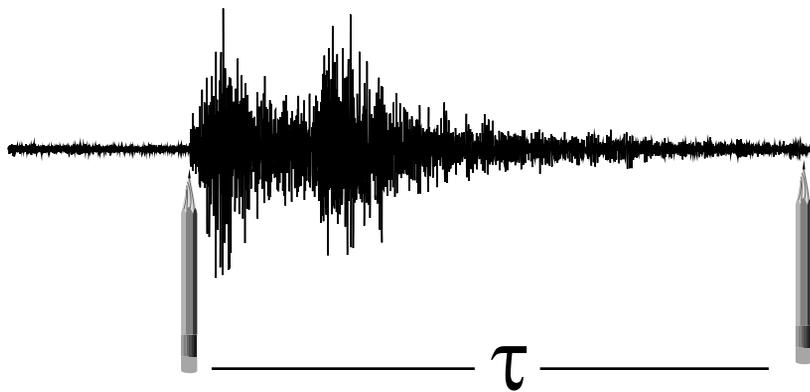


Figure 2 Signal duration (τ) as defined in this study with background noise on the portacorder records.

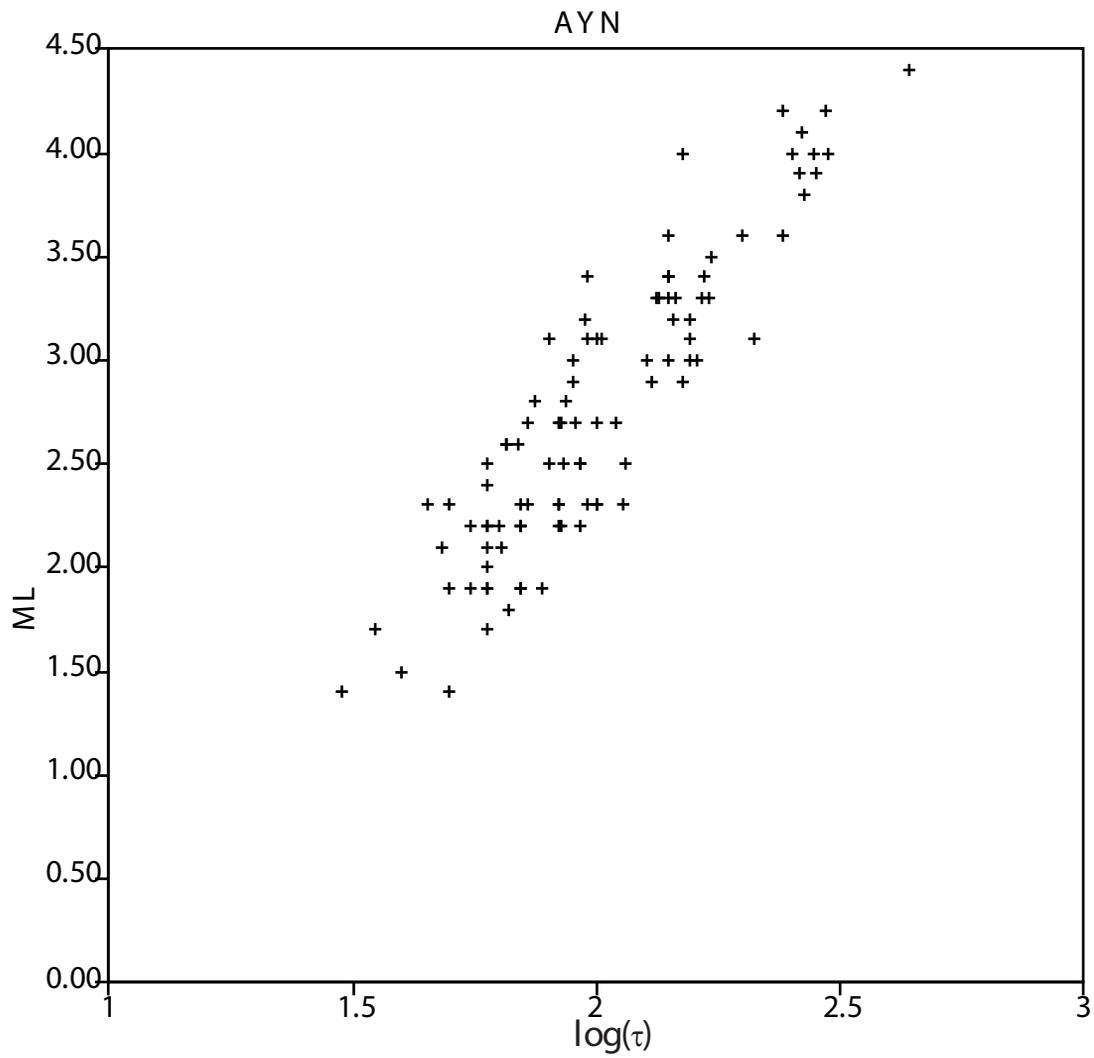


Figure 3 Correlation between ML and the common logarithm of the signal duration (τ) for 98 earthquakes recorded by AYN station.

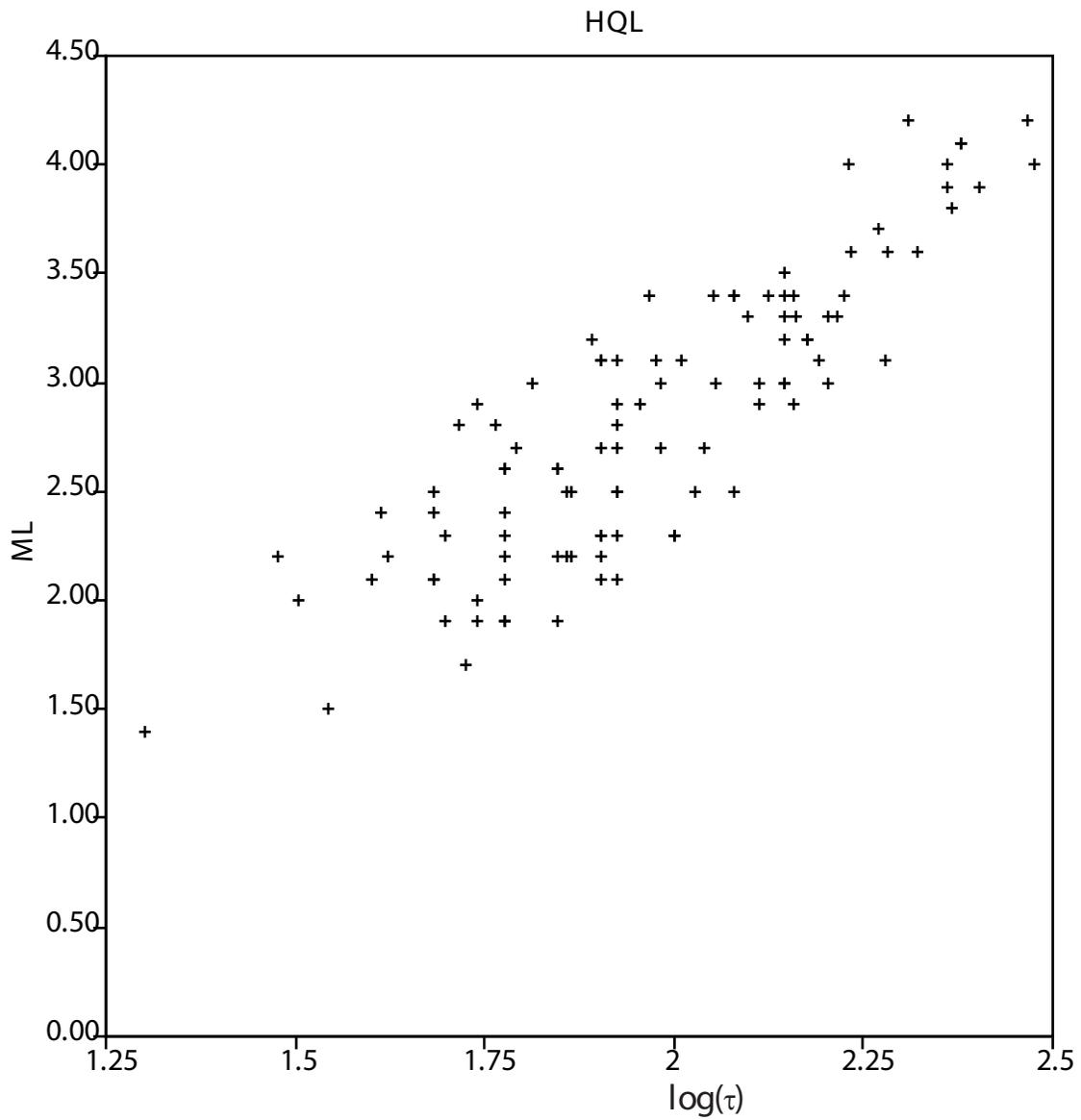


Figure 5 Correlation between ML and the common logarithm of the signal duration (τ) for 104 earthquakes recorded by HQL station.

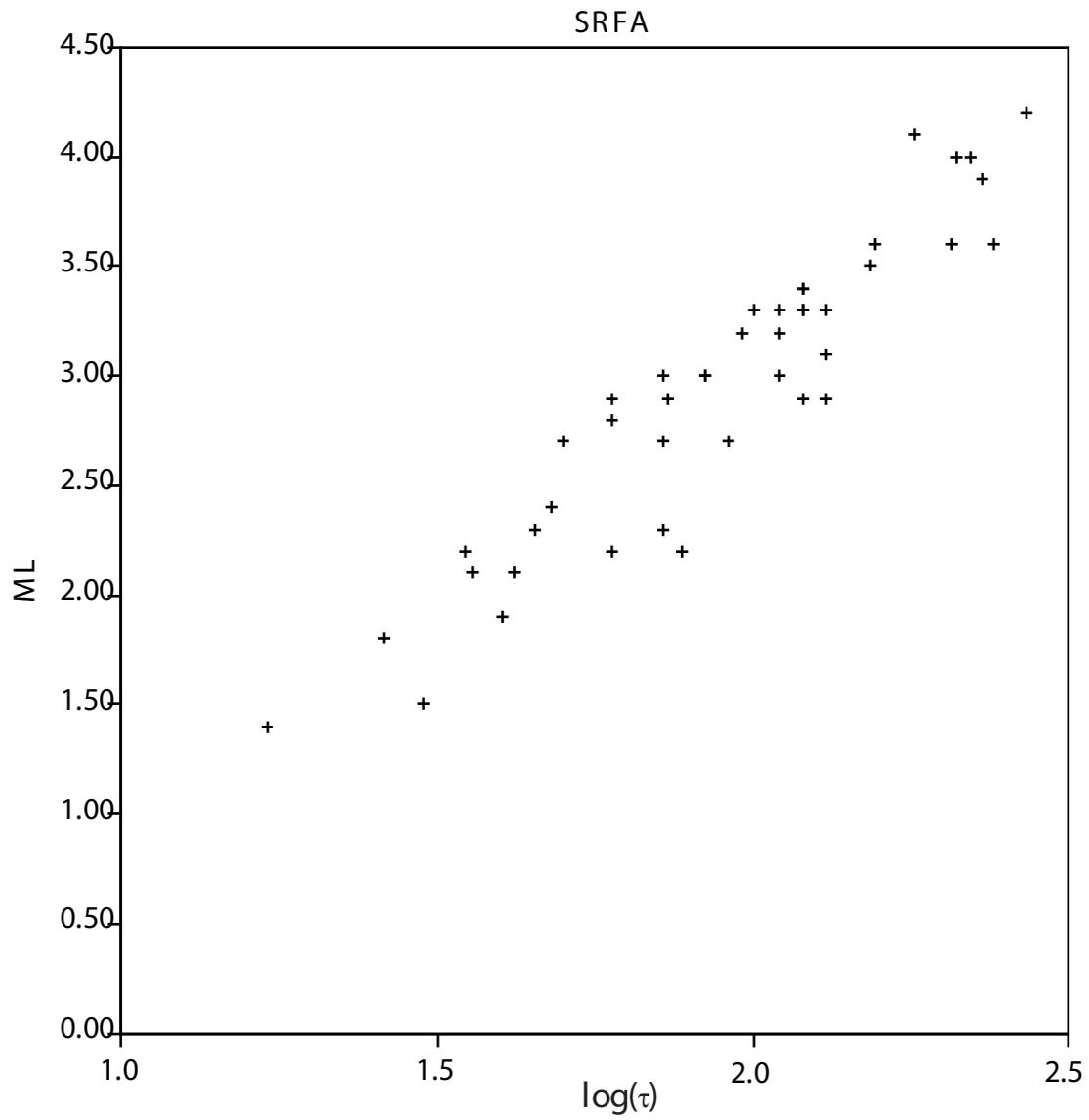


Figure 6 Correlation between ML and the common logarithm of the signal duration (τ) for 44 earthquakes recorded by SRFA station.